Hydrates in sediments and strategies for hydrate reservoir exploitation

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Gas Hydrates: A Major Energy Resource:





Hydrate Structures:



www.pet.hw.ac.uk/

Regions of pressure and temperature for stability of some hydrates

<u>Solid line</u> is liquid/ice coexistence curve

<u>Dashed line</u> is for methane hydrate

<u>Dash-dot</u> is for 2 per cent ethane in methane

+ are hydrate formed from water and an ethane-rich mixture with methane (3 per cent)



Where do the hydrate formers come from

- Biogenic sources
- Microbial activity in the upper several hundred meters of deep-sea sediment (about 99%)
- Thermogenic sources
- Thermal breakdown of organic material at greater depths



Seismic detection of hydrates

• Seismic methods: contrasts in density and elastic moduli associated with the presence of gas hydrate within sediment pore spaces versus saline pore fluid or free methane gas



Sampling of acustic reflection from seismic shots (from towed devices on ocean, close to bottom or bottom mounted acoustic signal generators)

Geophysical characterization and quantification of gas hydrates

P-cable: High resolution 3D seismic concept



- "Shallow" 3-D reflection seismic with high frequencies (50-250 Hz)
- A seismic cable towed perpendicular to the vessel's steaming direction
- Many seismic streamers attached to a wire (cross cable) held in place by two doors
- Used configuration with 12 (8) streamers towed parallel (streamer offset of ~10 m).

Geophysical evidence for gas hydrates

Seismic example from west Svalbard margin

Bottom-simulating reflector (BSR):

- Mimics seafloor
- Phased-reversed compared to seafloor reflection
- Reflection enhancement underneath





Buenz et al. (2003)

3D seismic system cube; from Vestnesa Ridge, W-Spitsbergen



Detection of hydrates through resistivity measurements

- Electrical methods are based on contrasts in electrical conductivity, or its inverse, resistivity.
- Gas hydrate is essentially an insulator, in contrast to saline pore fluids with resistivities on the order of 1 ohm-m or less

Source:http://marineemlab.ucsd.edu/ ~kkey/Pubs/Weitemeyer2006a.pdf



Example (Moridis & Collett, 2008) DOE Merit Review, August 26-27, 2008Pittsburgh, PENNSYLVANIA

- Resistivity typically estimates higher hydrate saturation than analysis from P-Wave velocity (seismic)
- (Even though these data appear to be open on the net the best is to request DOE for details on these data if interested in detailed location)



OUTLINE

- Before discussing exploration a few words about hydrates in sediments
- A few words about classical approaches for exploitation of hydrate reservoirs
- Simultaneous safe storage of CO2 and exploitation of hydrate reservoirs
- Theoretical modelling
- Conclusions

What is the present status of hydrate as a source of energy?

- The main worldwide motivations are essentially combinations of the following:
- 1) Desire for higher degree of energy supply independence (USA)
- 2) Limited own traditional oil and gas resources (Japan) or declining resources of conventional sources (India, South-East Asia)
- 3) Increased value of excisiting processing and transport infrastructure by connecting hydrate production to excisiting infrastructure.

So what is special about hydrates in sediments?

- Each different mineral has a specific structure and *surface structure of atoms*
- Interactions between these atoms and surrounding molecules does not facilitate stability of regular hydrate lattice in contact with the solid surfaces
- Exceptions are hydrates in contact with clay minerals, in which the ionic content of the clay liquids assist in "bridging" the hydrate to clay

Calcite cleaved along the dominant plane in contact with aqueous CO₂:



Calcite cleaved along the second most $st \{10\overline{10}\}$ plane in contact with aqueous CO₂: VMD-generated snapshot



Calcite cleaved along the second most st $\{10\overline{10}\}$ plane in contact with aqueous CO₂: density profiles



Example with a spherical hydrate particle growing from dissolved CO2 in saturated aqueous solutions at 1 C and 150 bar (x=0.033) x in minimum free energy for solution coexisting with the hydrate is x=0.016 System size is 40 nm x 40 nm and simulation time is 0.5 microseconds

- The chemical potential of the water molecules are approximated to vary linearly from the interface and assumed to be of bulk liquid properties 8 molecular diameters from the interface (based on structure sampling)
- Note that growth rate is highly sensitive to shape and a spherical particle has been chosen as a simple example
- A growing hydrate film shows similar behavior of "avoiding" the solid surface but the growth rate is slower



So what is the impact of this?

- Hydrate dissociation occurs wherever hydrate is exposed to *fluid channels undersaturated with respect to "guest"* types in the hydrate *and hydrate exposed towards undersaturated gas*
- I.e.: Hydrate phase transitions in porous media can occur on very large surfaces and channels between hydrate and minerals assists in mass transport
- For the opposite process it is enough to compare hydrate formation kinetics towards a liquid water surface as a reference (See later overheads for more detailed exp. setup)

Experiments with methane and water at 83 bar and 3 C Similar experiment with CO2 did not show any penetration of the hydrate film after 500 hour (resolution ~ 100 micrometer)



Note that methane is the wetting component on the polypropylene





Why? Two primary factors:

1) A methane hydrate film will rapidly form on the water/methane interface and reduce efficiently further growth untill film penetratesdue to local competition based on first and second laws of thermodynamics

2) Methane is the wetting component of the silicone rubber and some methane will migrate along the walls downwards in the chamber due

Heterogeneous hydrate formation (yellow) on the interface between CO2 (inner red) and liquid water initially saturated with CO2 (outer red)

Thermodynamics:		T = 274 K p = 150 bar 500 × 200 grid n = 100	
Fugacity coefficients for CO2 from SRK equation of state	Δx = 0.4 nm t = 0.27 μs		
Aquous description and hydrate description as in the homogeneous case	x _{CO2} =0.033	Pure CO ₂	x _{co2} =0.03 3
Simulations started with an extremely thin film of hydrate.			

Note that simulation times are very short for this simulation and the average growth rate at later stages are expected to be significantly smaller (as experienced for the homogeneous case)

Heterogeneous hydrate formation (red front) on the interface between CO2 (yellow) and liquid water (blue)

This simulation is veryy short and rates should be considered as initial rates and compared to theInitial rates from the homogeneous simulation the rate Is 7.5 times higher

DCO2=1.0-10-9 (liquid like) gives steady growth rate of 0.3 m/s

Experiment at T=277.4 K and 39 bar (lower thermodynamic driving force) Reported by Uchida et.al. (2003) indicates growth rate from 0.0001 m/s to 0.01 m/s



Similar simulation starting with a small hydrate nucleus

Average growth rate in this case is 0.24 m/s and diffusion-controlled. Note the short simulation time and recall that average rates for later growth are expected to be significantly lower



Now starting with several small hydrate nuclei on the interface between CO2 and water

Asymptotically the system behaves proportional to t in the exponent of 0.3 in contrast to 0.5 for total mass transport control







log(time)



Growth rates from dissolved CO2 much slower than heterogenous growth on exposed interfaces

So what is the consequence of this?

- Reservoir hydrate will never attach to the surface of the minerals
- The film of structured water, and possibly gas, will have a minimum thickness in the order of three water molecules but will normally be significantly thicker in a reservoir with significant fluid flow due to fractures and corresponding leakages

Consequence of mineral surfaces cont.?

- In practical terms this implies that real hydrate sections will have permeabillity, and a fluid filled porosity.
- But the magnitude of permeability may vary from very low to significant depending on the fluid flow dynamics of the reservoir (fractures, faults, feeding of fresh hydrocarbons from below etc)

Remarks

- Hydrate reservoirs which are tightly sealed (clay and shale layers) may enter a situation of extremely slow dynamics and local quasi equilibrium which will preserve the hydrate as "rich" in hydrocarbons. Example is Nankai Through
- Reservoirs which have symptoms of extensive "poc marks" of exposed hydrate may need more detailed examination since all leakage channels will have dissociated hydrate and the extent of rich and poor areas is uncertain



Sand dominated gas hydrate reservoirs





Hydrate Experiments Setup



A Bentheimer sandstone core, 4 cm diameter, 10 cm long is embedded into a heat controlled core holder

The hydrogen spin in liquid water and hydrate water is different. Exposing the system to magnetic field and sampling responses gives accurate (relative) feedback on amount of hydrate (invisible water hydrogen)

Permeability System

Additional measurements of pressure drop during hydrate formation gives possibility for recalculation of permeability



Time = 0 hrs



Time = 23 hrs


Time = 30 hrs



Time = 36 hrs



Permeability in Hydrates MRI Intensity Loss Indicates Hydrate Formation. Correlates with Permeability Decrease.



"Classical" approaches to hydrate production



Depressurization

Thermal Injection

Inhibitor Injection



Brings the system out of P,T stability zone but dissociation heat must be supplied from surroundings or added

Dissociation alone costs roughly 5% of produced HC Costly and require extra processing. Dilution from dissociated water Limits efficiency

Pressure and temperature are only two thermodynamic variable

Pressure reduction

Addition of heat



Inside or outside PT curve is just a minimum critera. Surrounding fluid concentrations and properties as well as mineral surfaces (adsorbed phase thermodynamics) are governing factors for hydrate phase transition dynamics The phase transition dynamics related to different exploitation schemes is a function of mass transport, heat transport and phase transition kinetics

 $\Delta H = -RT^2$

 The free energy difference is the driving force for the phase transitions kinetics (see next overheads)

 $\frac{\Delta G}{RT}$

 And the corresponding dissociation enthalpy must be transported from the surrounding at kinetic rates determined by the composite (mineral, hydrate, fluid) heat transport properties

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CO2 hydrate is more stable than methane hydrate

The free energy profile is constructed from MD simulations (Kvamme & Tanaka,1995) and verified to reproduce experimental stability regions (Svandal et.al., 2006)

Fig. 1 Free energy density of the hydrate as a function of the mole fractions of CH₄ and CO₂ at 1°C and 40 bars.



Fig. 2 Free energy density (kJ/mole) as a function of the mole fractions of CH_4 and CO_2 at 1C and 40 bars.

CO2 hydrate equilibrium pressures

Squares :experimental data Solid line: predicted results



References:

Svandal, A., Kuznetsova, T., Kvamme, B., Physical **Chemistry Chemical** Physics, 2006, 8, 1707 - 1713 Svandal, A., Kuznetsova, T., Kvamme, B.,, 2006, Fluid Phase Equilibria, 2006, 246, 177-184 Kvamme, B. & Tanaka, H.,1995, J.Phys.Chem., 99,

7114

Before discussing exploitation schemes we need to look at the most relevant situations

- Type 1 hydrate occurences
 Hydrate on top of liquid water and a gas cap facing the hydrate
- Type 2 hydrate occurences
 Hydrate on top og liquid water with no free gas
- Type 3 hydrate occurences
 Solid hydrate without any contacting fluid

Type 2 and 3 hydrates

- Hydrate is generally not stable toward liquid water undersaturated with hydrate former (realtive to free energy of hydrate!)
- The consequence is that hydrate sections might be interrupted by liquid water sections due to substantial fluid flow (dilution of fluids surrounding hydrate structures) in these regions
- Type 3 hydrates are more common in systems where shale or clay layers separate hydrate blocks

Some very general factors

- Stability of hydrates is very sensitive to temperature – which practically implies that efficiency of hydrate exploitation increases with higher in situ temperature.
- Production potential increases with intrinsic and relative permeability

Type 1 hydrates and *pressure reduction approach*

- Pressure reduction is
 feasible but slow
- Efficiency can be improved through local fracturing
- Kinetics of hydrate dissociation likely to be dominated by heat transport limitation from surroundings in order to supply the dissociation enthalpy needed

Knowledge gaps and limitations

- Heat transport dynamics is fairly complex and no satifactory models available
- Depressurisation reduces temperature and may lead to refereezing and in the worst case also ice formation

Knowledge gaps and limitations continued

- It remains unverified how much of the released water that will be transported with the gas
- Most efficient exploitation from uncosolidated sediments with significant porosity but this type of sediments are also more sensitive to potential geomechanical implications of the volume reduction (~ 10%) by dissociation of hydrate
- Efficiency depends sensitively on types of producuing well(s) (horizontal or vertical)

Addition of heat?

- Some limited additional heat is necessary anyway to keep producing lines free of possible refrozen hydrate and/or ice.
- Injection of heat through steam or brine is possible but involves heat losses beyond the 5% of reduced hydrocarbon value involved in the dissociation alone.
- System of heat injection wells and possible local fracturing determines efficiency together with heat transport properties of the total system

Injection of inhibitor?

 Injection of inhibitor can also be used as alternative to heat in order to keep production channels open but then again requires corresponding separation facilites of separating inhibitor from produced water.

Summary

- Type I hydrate structures are fairly easy to produce by means of pressure reduction, which will simultaneously produce the free gas below
- And except from strategies for preventing reformation of hydrate and/or ice formation the system does not require any new or special technology
- But geomechanical considerations are needed since primary targets are unconsolidated sand hydrate systems

Type 2 hydrate

- Pressure reduction with/without addition of heat is possible but significant production of water can not be avoided.
- Efficiency still unverified due to limitations in description of hydrate dissociation dynamics.
- Horisontal producing wells will be an efficient completion and no special technology is needed for producing these hydrates

Type 2 hydrate cont.

- Dissociation rates will be low but can be increased by local fracturing as well as chemical injection
- This type of situation is also an ideal target for exploitation though CO2 injection, as will be explained later

Type 3 hydrate

- As discussed earlier there is no unique type 3 hydrate.
- It all depends on the actual permeability due to the state of dynamics ("tight" reservoirs with little fluid flow to dynamic reservoirs with substantial fluid flow)
- Dynamic reservoirs with significant permeability can be produced by pressure reduction with the aid of some local fracturing to increase exposed hydrate surfaces and increase average permeability

Alternatives?

- Chemical injection is possible for type 2 and all ranges of type 3 hydrates but is not likely to be economically feasible
- Injection of carbon dioxide
- Injection of flue gas





exploitation of hydrate reservoirs?

Storage of

CO2

and

Can it be experimentally verified that this will practically work in porous medium?

Project Objective

Experimentally Verify:

- Sequestering Greenhouse Gas (CO₂)
 Determine hydrate formation and distribution
- Gas Production from Hydrates

 Determine the rate of CO₂-CH₄ exchange

Numerically Predict:

- The hydrate reformation and gas production

Experimental Approach

Use Magnetic Resonance Imaging (MRI) to Image the Formation and Dissociation of Gas Hydrates In Porous Media.

Hydrate Experiments Setup



Experimental Setup



Experimental Setup



Core Sample Design

- Bentheim Sandstone Water-Saturated
- Longitudinal Cut Fitted Spacer Simulates Open Fracture
- Experimental Conditions: Flow Loop ~ 4°C – 8.3 MPa (1200 psi).
























































1 2

Time – 214hrs



Methane starts to fill fracture

Time – 230hrs



Methane starts to fill fracture







Methane starts to fill fracture















Time – 327hrs







Time – 483hrs









CH₄ Production Rates & Amounts from Hydrate



Theoretical modelling

- The conversion of CH4 hydrate over to CO2 hydrate releases limited amounts of energy
- Heat transport is, however, 2 orders of magnitude or more faster than mass transport.
- As a first approximation we may treat this particular exchange as isothermal

Phase Field Theory

- Density Functional Theory links the kinetics of phase transition to the change in molecular structure across the phase boundaries
- Molecular structures is uniquely linked to corresponding free energies via statistical mechanics
- Phase Field Theory uses free energy changes directly as the driving forces for kinetic progress of the phase transition

Three component Phase Field Theory

$$F = \int d\underline{r} \left\{ \frac{\varepsilon^2 T}{2} (\nabla \varphi)^2 + \sum_{i,j=1}^3 \frac{\varepsilon_{i,j}^2 T}{4} (c_i \nabla c_j - c_j \nabla c_i)^2 + f_{bulk}(\varphi, c_1, c_2, c_3, T) \right\}$$

$$f_{bulk} = wTg(\varphi) + [1 - p(\varphi)] f_s(c_1, c_2, c_3, T) + p(\varphi) f_L(c_1, c_2, c_3, T)$$

$$\dot{\varphi} = -M_{\varphi} \frac{\delta F}{\delta \varphi} + \zeta_{\varphi}$$

$$\sum_{i=1}^3 c_i = 1$$

$$\dot{c}_i = \nabla M_{ci}(c_1, c_2, c_3) \nabla \left(\frac{\delta F}{\delta c_i} - \zeta_i\right)$$

Parameters ϵ and w can be fixed from the interface thickness and interface free energy. ϵ ij set equal to ϵ



Fig. 6 Full line -- hydrogen density profile; dashed line -- oxygen in CO₂ dashed-dotted line -- oxygen in water (both hydrate and liquid).
TABLE I. 10%-90% interface widths from MD simulations

	d (nm)	×
H₂O D↑	0.93	z
H₂O D↓	0.86	had a server a water to the children book to
CO₂ D↑	0.79	
H₂O CD↑	0.61	
$H_2O CD\downarrow$	0.69	
$CO_2 CD^{\uparrow}$	1.21	
$CO_2 CD\downarrow$	0.88	
Average	0.85	
H ₂ O DGF	1.21	
CO ₂ DGF	1.07	

Notation: D – mass density, CD – charge density, \uparrow – envelope of peaks, \downarrow – envelope of wells, GF – Gaussian filtered.

Cleaving the water -- CO₂ hydrate system by adding repulsive interactions:



Error bars too large and enhanced by tendencies of hydrate dissociation at the interface. Work on alternative approach according to cappilary wave theory is in progress

Interface thickness d and interface free energy fixes the two model parameters *w* and ε in the Phase FieldTheory. Evaluation of these parameters from model systems by Molecular Dynamics simulations of model systems

Estimated to 8.5 Å – see previous overhead

$$d = \left(\frac{\varepsilon^2 T}{2}\right)^{1/2} \int_{0.05}^{0.95} d\xi \left\{ \Delta f[\xi, c(\xi)] \right\}^{-1/2}$$

Phase Field Theory parameter w is in the free energy f (see PFT slide)

$$\gamma_{\infty} = (\varepsilon^{2}T)^{1/2} \int_{0}^{1} d\xi \{\Delta f[\xi, c(\xi)]\}^{1/2}$$

Phase Field Theory parameter ε

Interface free energy ≈ 30 mJ/m²

where
$$\Delta f = f - f_0$$
, and

$$f_0 = f_L(c_L^{eq}) + \left(\frac{\partial f_L}{\partial c}\right)\Big|_{c_L^{eq}}(c - c_L^{eq})$$

$$= f_{S}(c_{S}^{eq}) + \left(\frac{\partial f_{S}}{\partial c}\right)\Big|_{c_{S}^{eq}}(c - c_{S}^{eq})$$



Fig. 2 Free energy density (kJ/mole) as a function of the mole fractions of CH_4 and CO_2 at 1C and 40 bars.



Fig. 1 Free energy density of the hydrate as a function of the mole fractions of CH_4 and CO_2 at 1°C and 40 bars.

Phase field predictions for the conversion of a 10 nm thick methane hydrate layer into mixed CO2 – CH4 hydrate in the presence of liquid CO2 at T =276.15 K and p = 8.3 MPa. The liquid CO2 phase is on the right. The spatial step size is $\Delta x = 0.05$ nm.



- (a) the phase field (solid), CO2 (dashed), CH4 (dotted), and H2O (dash-dot) concentration profiles corresponding to instances t = 0.0, 0.08, 0.16, 0.24, 0.32, and 0.4 μ s. The initial profiles are shown by thinner lines. Note the thin water layer slightly right of the z = 10 nm position, and the respective depletion of methane and carbon-dioxide in the same region.
- (b) An enlarged view of the methane and carbon-dioxide profiles in the solid hydrate phase is presented in panel (b).

Detailed analysis of the simulated results indicate that the conversion process is essentially dominated by transport limitations. I.e.: we should expect that Fick's law would represent the kinetic progress of conversion to a fair extent



Displacement *h* of the half-height ($c_{CO2} = 0.055$) position for the CO₂ profile as a function of $t^{1/2}$. $D_s = 5 \times 10^{-12}$ m²/s is assumed.

So how does this value of diffusivity match experimental results



"Bulk" experiment without porous medium indicate D ~ $1.1 \cdot 10^{-11}$

Effects of liquid channels separating the hydrate from the mineral surfaces and the state of non-equilibrium is systems with mineral surfaces (adsorption is an additional phase which reduces the number of degrees of freedom with 1) may account for the difference





Conclusions

- All hydrate reservoirs do have permeability for the simple reason that water hydrogen bonding in hydrate does not match any favorable interactions with mineral surfaces
- Hydrate phase transition dynamics and corresponding producability depends on type of situation (underlying gas cap, underlaying water or hydrate block alone) depends on the flow dynamics of the actual reservoir in consideration

Conclusions cont.

- Type I reservoirs (undelying gas cap) can be produced initially by pressure reduction but as produced gas is replaced by groundwater it will gradually turn into type II, which is very suitable for production through CO2 injection.
- Local fracturing and horisontal wells for injection of CO2 between water and hydrate will increase injectivity and at the same time create escape pathways for methane

Conclusions cont.

- We have demonstrated experimentally that injection of CO2 into CH4 hydrate results in CH4 release while at the same time providing safe storage option for CO2
- Unlike other concepts for releasing the hydrocarbons from hydrate there is not net volume change, which can be a geomechanical concern when dissociating hydrate

Conclusions continued

- No extra production of liquid water associated with the gas exchange
- No additional energy supply needed (the conversion is actually exothermic)
- Visualizations Provide Different and Helpful Insights of Hydrate Growth and Stability
- Theoretical predictions based on molecular simulations and Phase Field Theory reproduce the experimental conversion rates

Fiery Ice from the Seas



A series of workshops arranged by

Bjørn Kvamme, UoB, Norway Stephen Masutani, UoH, USA Rick Coffin, ONR, USA, Tsutomu Uchida, Hokkaido Univ., Japan

7th workshop May 10 – 12, 2008 **Te Papa, Wellington, New Zealand** http://www.gns.cri.nz/fieryice/ **Organized by: Ingo Pecher**

Purpose: Contribute to the establishment of large international interdisciplinary research programs on fundamental properties of hydrate reservoirs, hydrate quantification and feasible exlotitation